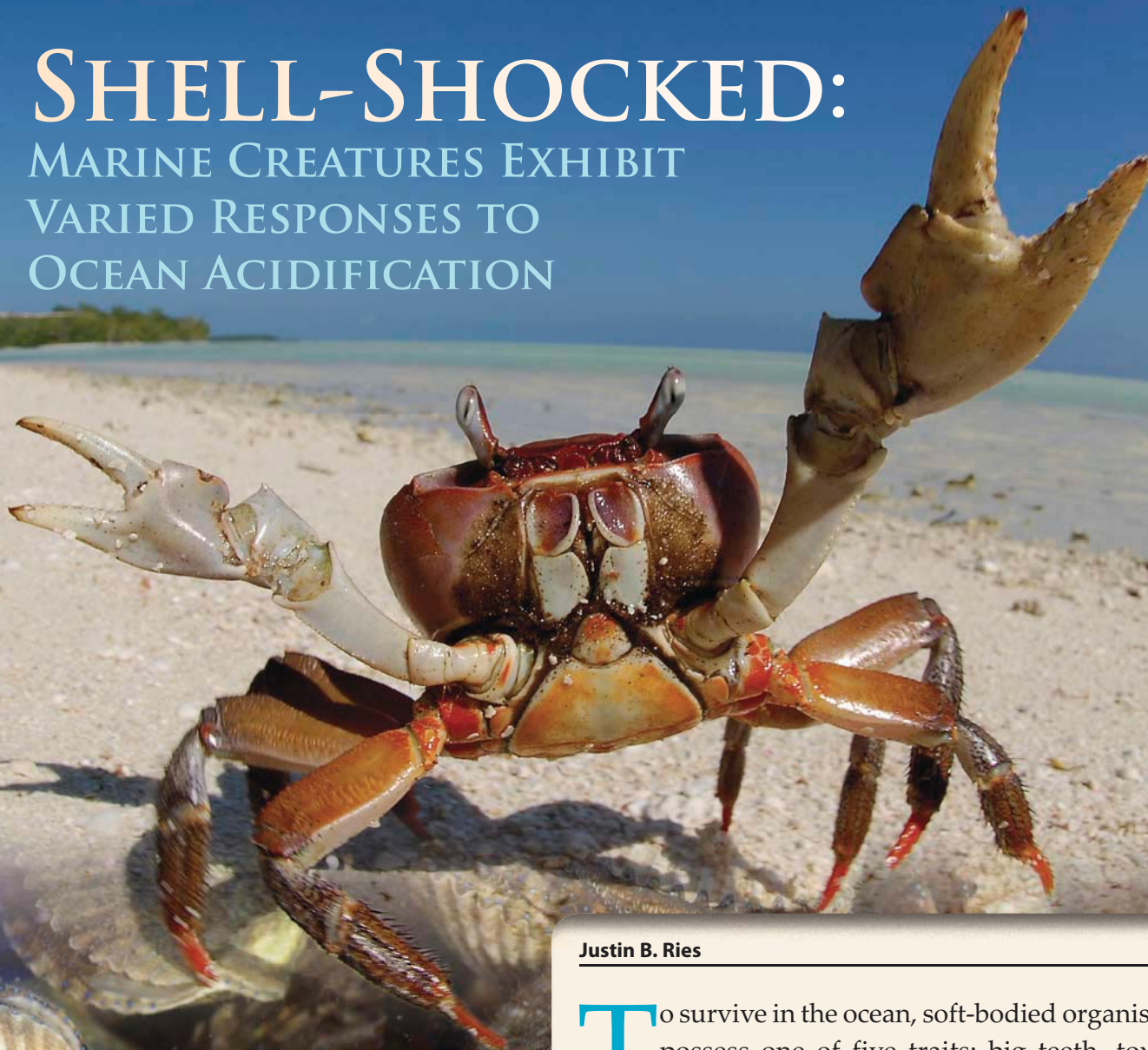


# SHELL-SHOCKED:

MARINE CREATURES EXHIBIT  
VARIED RESPONSES TO  
OCEAN ACIDIFICATION



Justin B. Ries

**T**o survive in the ocean, soft-bodied organisms must possess one of five traits: big teeth, toxic flesh, invisibility, quickness or a hard shell. Most marine organisms that employ the latter, called calcifiers, build their hard shells from the mineral calcium carbonate. However, increasing atmospheric carbon dioxide levels are making the oceans more acidic — which, in turn, is reducing the concentration of carbonate ions dissolved in seawater that organisms use to build their protective shells and skeletons. Calcifying marine organisms — such as clams, crabs, corals and conchs — may soon find themselves short on building material.

However, a recent set of experiments suggests that the ocean acidification story is more complex than first thought. Whereas some marine calcifiers reared under the elevated carbon dioxide levels responded very negatively, not all of the organisms suffered in the acidified seawater. Some, in fact, appeared to benefit from it.

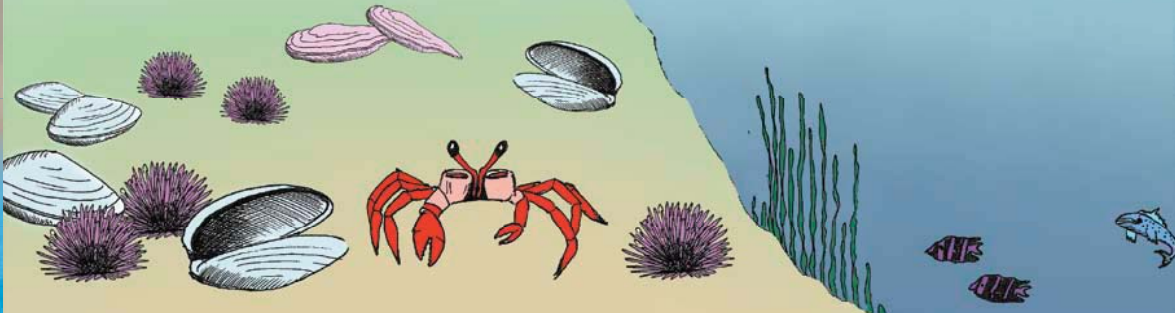
Bottom: ©iStockphoto.com/ShaneKato; top: ©iStockphoto.com/Rainer von Brandis

more carbon dioxide emissions from cars, powerplants,  
cement factories and deforestation



more carbon dioxide + water => more carbonic acid => more bicarbonate + more hydrogen ions

more hydrogen ions + carbonate => more bicarbonate



calcium + less carbonate => less calcium carbonate for shells

As atmospheric carbon dioxide levels increase, the oceans are absorbing much of the additional carbon dioxide, forming carbonic acid in the seawater. That increases the acidity of the waters (adding more hydrogen ions). However, the oceans also contain a natural pH buffer — dissolved carbonate and bicarbonate ions. Those ions react with the additional hydrogen ions in the water, partially neutralizing the change in acidity. But those reactions also leave fewer carbonate ions available for calcifying organisms to build their shells — and fewer dissolved carbonate ions (a lower “saturation state”) can also make the shells themselves more vulnerable to dissolution.

## SETTING THE SCENE


Carbon dioxide has fluctuated drastically throughout Earth’s past. Since the evolution of primitive marine calcifiers about 550 million years ago, atmospheric carbon dioxide is thought to have fluctuated between approximately 200 and 8,000 parts per million (ppm), with highest levels occurring in the Early to Middle Paleozoic (about 543 million to 400 million years ago), Late Triassic/Early Jurassic (about 200 million years ago), and Cretaceous (about 125 million years ago). Carbon dioxide levels dropped after the Cretaceous, however, and have not been as high as today in the past 800,000 years.

Carbon dioxide levels in our atmosphere are currently at about 385 ppm. They have been rising steadily since the late 1700s, due in large part to the burning of fossil fuels and deforestation. Modelers predict that if we don’t act to stem carbon

dioxide emissions in the near future, atmospheric carbon dioxide levels could reach 750 ppm by the year 2100.

Given the extreme fluctuations in atmospheric carbon dioxide that are thought to have occurred throughout Phanerozoic time (the past 543 million years), skeptics of the ocean acidification threat often argue that many types of marine calcifiers have already survived carbon dioxide levels that dwarf those predicted for the coming centuries. Indeed, the high-carbon dioxide Cretaceous period is even named for the massive chalk formations formed from the shells of tiny calcareous plankton that flourished during this interval.

The truth of the matter is that we don’t know exactly what will happen to calcium carbonate-producing organisms when carbon dioxide levels hit certain thresholds, especially when these changes are not accompanied by the elevated seawater



alkalinity that probably accompanied intervals of high atmospheric carbon dioxide in the geologic past. Thus, we set out to simulate future acidic oceans and to investigate the responses of marine calcifiers for ourselves.

### CLAMS GET CRUSHED, CRUSTACEANS GET CRUSTY

We tested 18 species of marine calcifiers under four different atmospheric carbon dioxide scenarios, ranging from the current level (roughly 400 ppm) up to 2,850 ppm.

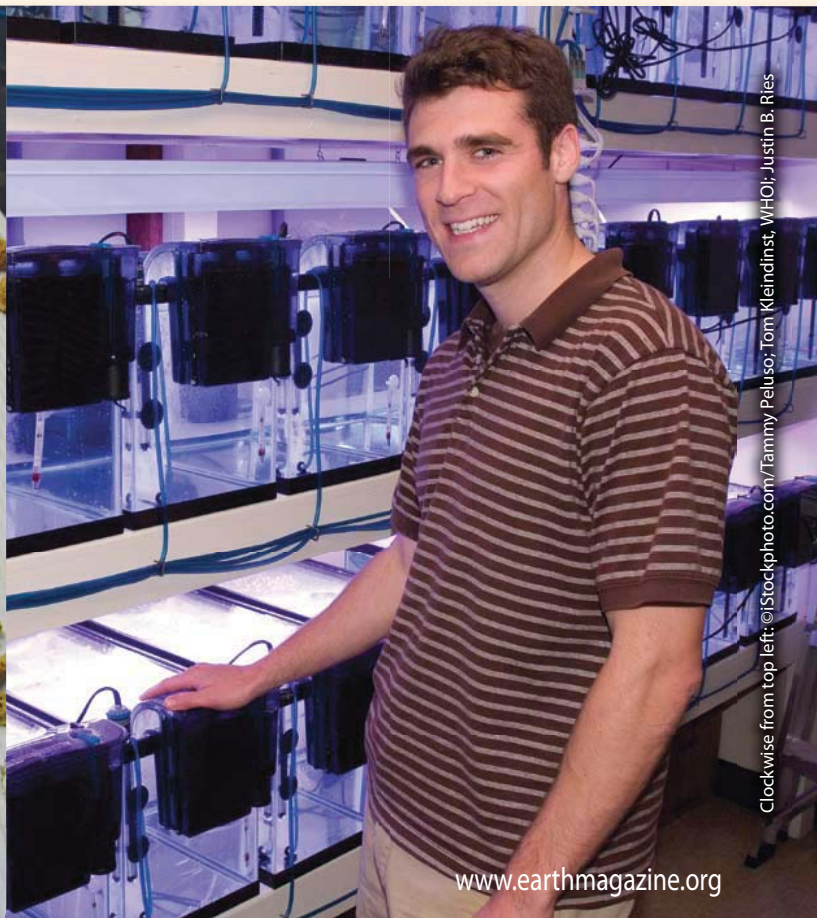
Mollusks — including oysters, quahogs, softshell clams, bay scallops and conchs — fared by far the worst under the elevated carbon dioxide scenarios. This is bad news not only for shellfish lovers, but also for the local, state and federal governments that reap substantial tax receipts from the billion-dollar industries based upon these briny delicacies.

The mollusks' responses varied widely amongst species, however. Bay scallops, periwinkles, whelks, oysters and softshell clams built their shells more and more slowly as carbon dioxide increased. The conchs and the quahogs showed no response to carbon dioxide levels up to 900 ppm — but above 900 ppm, they showed a very negative response. One type of mollusk, the slipper limpet, showed a particularly surprising response. Its calcification rate actually increased under rising carbon dioxide levels up to 900 ppm, and only showed a decline under the highest carbon dioxide treatment (2,850 ppm). Intriguingly, the tasty blue mussel did not respond at all to the elevated carbon dioxide.

We also tested two types of urchins — tropical pencil urchins and temperate purple urchins — and two species of calcifying algae, the branching coralline red algae *Neogoniolithon* and the segmented calcareous green algae *Halimeda*. Tropical urchins

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Below left: Justin Ries looked at 18 species of ecologically and economically important marine calcifiers. The species investigated represent eight of the major marine calcifying groups, including crustaceans, corals, bivalves, gastropods, coralline red algae, calcareous green algae, serpulid worms and urchins. The investigated species produce their shells and skeletons from the aragonite, high-magnesium calcite and/or low-magnesium calcite forms of calcium carbonate. Below right: Ries and his experimental ocean acidification system. He and his colleagues raised marine calcifiers for 60 days in experimental seawaters bubbled with mixed gases formulated at carbon dioxide partial pressures of 400, 600, 900 and 2,850 parts per million (ppm) to determine how such atmospheric carbon dioxide levels would affect the marine creatures.



Clockwise from top left: ©Stockphoto.com/Tammy Peluso; Tom Kleindinst, WHOI; Justin B. Ries

As atmospheric carbon dioxide levels increase, marine calcifiers react differently. Some, such as bay scallops, fared very poorly under elevated carbon dioxide levels, whereas others, such as blue crabs, fared well. The new research measured the response of marine organisms to 400, 600, 900 and 2,850 ppm atmospheric carbon dioxide (increasing to the right of the graph).

showed no response between 400 and 900 ppm carbon dioxide, and then began to dissolve away rapidly at 2,850 ppm. Temperate urchins and both species of algae exhibited increased calcification up to 900 ppm carbon dioxide followed by a decline at 2,850 ppm.

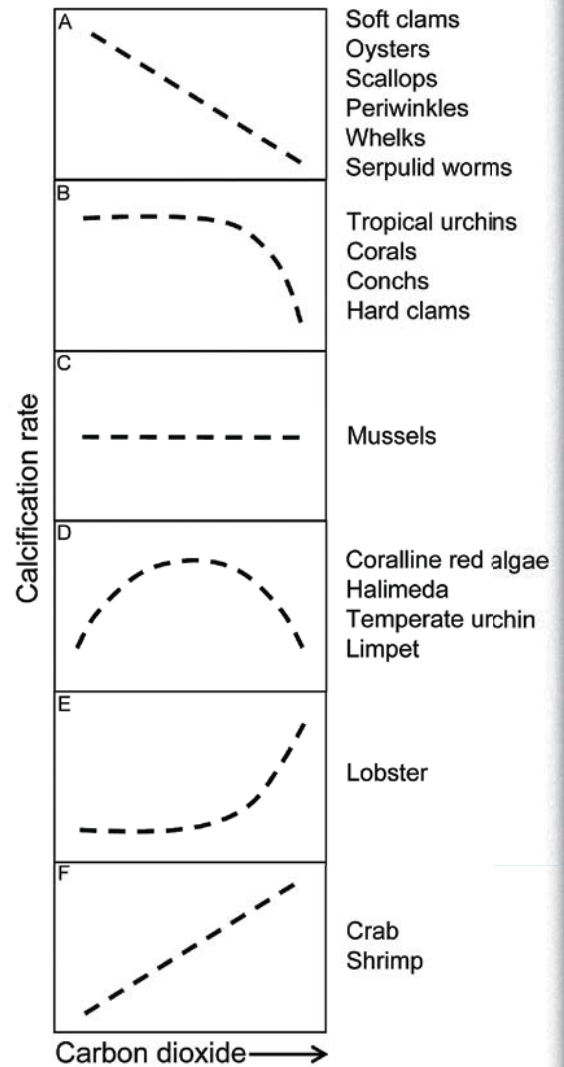
Yet it was the three species of crustaceans — blue crabs, gulf shrimp and American lobsters — that exhibited the most striking response of all. They each calcified most rapidly under the highest carbon dioxide level. The response of blue crabs and the shrimp was fairly linear, with calcification rates increasing steadily with rising carbon dioxide. The lobsters showed no response to elevated carbon dioxide between 400 and 900 ppm, and exhibited an increase in calcification under the highest level.

Although not anticipated, these varied responses should not be surprising. Many of the organisms investigated in these experiments belong to groups that survived intervals of the geologic past in which atmospheric carbon dioxide was 10 to 20 times greater than that of today. Thus, these marine calcifiers must have evolved strategies for coping with such acidified conditions. And indeed, these organisms are thought to employ a wide range of relatively sophisticated mechanisms for building their shells and skeletons. So what factors are at play?

First, most calcifying marine organisms cover their shell or skeleton with some type of protective organic layer that separates them from ambient seawater. In general, organisms that produce a relatively thick organic layer that covers most or all of their shell or skeleton, such as the crustaceans, the algae and the blue mussels, were generally more resilient to elevated carbon dioxide than organisms that produce a less substantial protective barrier, such as the clams, oysters, scallops and conchs. Surely this barrier comes at a metabolic cost to the organism, but is apparently well-worth the investment when ambient conditions become too acidic.

It has been proposed that the type of calcium carbonate secreted by the organism — aragonite, low-magnesium calcite or high-magnesium calcite — should play a large role in determining an organism's specific response to carbon dioxide-induced ocean acidification. This is because the aragonite and high-magnesium calcite forms of calcium carbonate are more soluble than the low-magnesium calcite form. But the experiments revealed that the role of mineralogy is

A thicket of staghorn corals on the Meso-American Barrier Reef System off of Belize. Scleractinian corals such as these build their skeletons from the more-easily dissolved aragonite form of calcium carbonate. Because of the relatively high solubility of aragonite, these corals may be particularly vulnerable to carbon dioxide-induced ocean acidification.





The aragonite shells of conchs reared under current (left, 400 ppm) and elevated (right, 2,850 ppm) atmospheric carbon dioxide levels. The conch shell reared under elevated carbon dioxide exhibits clear signs of dissolution. The high-surface-area knobs along the upper lip of its spiral shell have largely dissolved away. Conchs may have evolved these knobs for stability on the seafloor.

not so straightforward. Crustaceans, for example, secrete a high-magnesium calcite shell (8 to 12 percent magnesium) but exhibited a more positive response to elevated carbon dioxide than all other investigated organisms, including those secreting low-magnesium calcite, like oysters, bay scallops and periwinkles.

Instead, it appears that the blue crab's epicuticle layer — the outermost part of the exoskeleton — protects its relatively soluble high-magnesium calcite exoskeleton from the corrosive seawater. Shell mineralogy did appear to play a predictable role under the highest carbon dioxide level, however, as five of the six species that dissolved under these conditions formed their shells from high-magnesium calcite and/or from aragonite — forms of calcium carbonate that are less stable than calcite.

Carbonate is the most important form of dissolved inorganic carbon utilized by marine organisms in the calcification process. Yet another form of dissolved carbon — bicarbonate (carbonate plus a proton) — is 10 times more abundant than carbonate in seawater and, unlike carbonate, actually increases in seawater with rising atmospheric

carbon dioxide. Organisms cannot directly use bicarbonate in calcification, but some are believed to be able to convert it to carbonate for use. Crustaceans and calcifying algae are thought to be particularly adept at this process. Thus, it is perhaps no coincidence that these organisms exhibited relatively positive calcification responses to elevated carbon dioxide.

### FERTILIZING PHOTOSYNTHESIS

We also investigated how photosynthesizing marine calcifiers would react to high carbon dioxide conditions. In these organisms, photosynthesis and calcification have long been viewed as complementary processes: Photosynthesis removes carbon dioxide from seawater, changing the seawater's saturation state — less dissolved carbon dioxide means more dissolved carbonate ions. Likewise, calcification releases carbon dioxide that the organism needs for photosynthesis.

Increased dissolved carbon dioxide in seawater (resulting from increased carbon dioxide in the atmosphere) may also promote more photosynthesis, which could provide the organism with additional energy for calcification. This would result in enhanced calcification under conditions of elevated

## ON GROWTH AND FORM

**M**arine shells are some of the most beautiful and complex structures on Earth. Yet nearly every spine, ridge and curve within these shells has evolved to fulfill some function.

For example, conchs are thought to have evolved bumps on their spiraled surface to prevent them from spinning in endless circles on the seafloor as they are swept by the

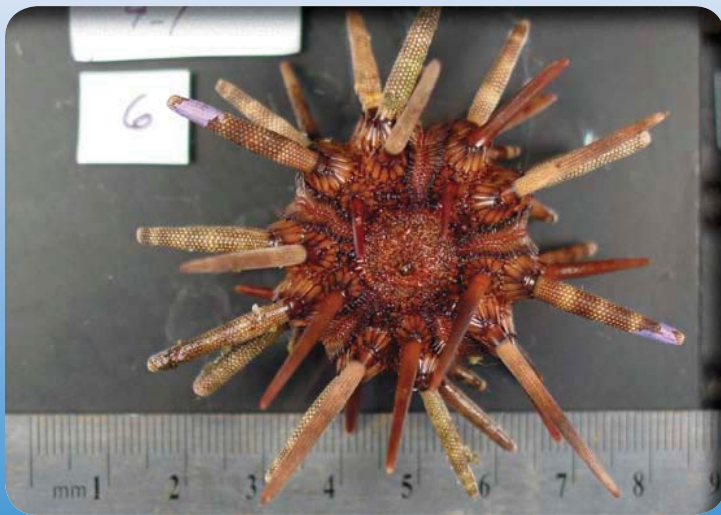
current. Similarly, clams are thought to have evolved asymmetrical ridges along their shells' outer surface to help them grip and wiggle down through the sediment.

Conchs reared under the highest carbon dioxide treatment in our experiments began to lose their shell bumps through dissolution. And the shell ridges on the quahogs reared under the same conditions became

smooth and flat. The tropical urchin's spines, used for protection and motility, also began to dissolve away.

The carbon dioxide-acidified seawater likely targets these functional features because of their relatively high surface area. More experiments need to be run to test the full impact that ocean acidification will have on the functional morphology of marine calcifiers.

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Tropical pencil urchins (high-magnesium calcite) reared under current (left, 400 ppm) and elevated (right, 2,850 ppm) atmospheric carbon dioxide. The spines of the pencil urchins are well-formed under current carbon dioxide conditions, but are largely dissolved away under the highest carbon dioxide treatment. Urchins use their spines for protection and motility.

carbon dioxide despite a reduction in the calcium carbonate saturation state of the surrounding seawater.

Even though the three photosynthesizing organisms (two algae species and a coral species) we investigated exhibited either a positive or neutral response to elevations in carbon dioxide between 400 and 900 ppm, all three exhibited a substantial reduction in calcification under the highest carbon dioxide level.

This suggests that enhanced calcification due to the additional photosynthesis occurs only when carbon dioxide is the limiting factor for the rate of photosynthesis. Once there is more than enough carbon dioxide for these organisms — at around 1,000 ppm for marine algae — then the benefits of increased photosynthesis would be quickly offset by the negative effects of having fewer carbonate ions in the surrounding seawater.

### MARINE CALCIFIERS IN HOT WATER?

Determining whether warm-water or cold-water organisms are more vulnerable to rising levels of atmospheric carbon dioxide is a pressing goal of ocean acidification research. Carbon dioxide is more soluble in cold water than in warm water — so colder waters tend to be more acidic. Organisms at higher latitudes and inhabiting deeper, colder waters are therefore expected to fare worse than organisms living in the tropical seas.

However, it seems equally plausible that the organisms living in those colder, more acidic waters would have already evolved to be more resistant to such changes in the carbonate chemistry of seawater. Indeed, the tropical urchin we tested was more negatively affected by the acidic seawater than the temperate urchin was. This may also explain why the temperate coral we





A close-up photograph of a blue crab resting on a piece of weathered, greyish-brown wood. The crab's body is a mix of green and blue, with its legs and claws showing vibrant blue and red colors. The wood has a vertical grain and shows signs of being outdoors, with some knots and texture visible.

## A CARBON SINK ON THE BRINK

**T**he largest carbon sink on Earth is carbonate rock — limestone and dolostone, in particular. These rocks are formed primarily from the skeletons of marine calcifiers.

Our experiments suggest that calcareous marine organisms that cover their outer shells and skeletons with protective organic layers are able to maintain positive rates of calcification even under extremely elevated carbon dioxide conditions. However, the functionality of these organic layers will inevitably decline after the death of these organisms, resulting in rapid post-mortem dissolution of their shell under high-carbon dioxide conditions.

This could effectively shut down Earth's largest carbon sink, thereby triggering a positive feedback effect within the global carbon cycle that could accelerate the rate at which carbon dioxide builds up in the atmosphere.

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interval have also recently been attributed to global carbon dioxide-induced ocean acidification events.

Secondly, it is not simply atmospheric carbon dioxide that determines the saturation state of seawater with respect to calcium carbonate minerals. Many factors play a role, not least of which are seawater alkalinity (a property related to the concentrations of carbonate and bicarbonate ions in seawater), temperature and the other (little-mentioned) ion involved in shell building: calcium.

Calcium is viewed as a conservative element in modern seawater: Its concentration relative to other ions dissolved in seawater remains relatively constant throughout the oceans. However, geochemical evidence suggests that the calcium concentration of seawater has fluctuated markedly throughout geologic time, in large part due to changes in the flux of brines from hydrothermal vents along mid-ocean ridges and large igneous provinces. In short, high rates of ocean crust production result in elevated concentrations of calcium in seawater.

On the other hand, ocean crust production and the resulting volcanism are also thought to be important contributors of atmospheric carbon dioxide over geologic time. Therefore, throughout Phanerozoic time, these tectonic drivers of seawater chemistry may have partially offset one another in terms of their net

effect on shell-building: Increased ocean crust production would reduce the carbonate ion concentration of seawater (due to elevated atmospheric carbon dioxide), yet would simultaneously increase its calcium concentration. This suggests that marine calcifiers may not have experienced such extreme fluctuations in the calcium carbonate saturation state of seawater, after all.

Lastly, if tectonically driven fluctuations in atmospheric carbon dioxide in the geologic past occurred more gradually than the current anthropogenic rise in atmospheric carbon dioxide, then acidified seawater, as well as acidified rainwater, would have increased the dissolution of carbonates on the ocean floor and on continents. That would have increased the carbonate alkalinity of seawater, which may have rendered elevations in atmospheric carbon dioxide less harmful to marine calcifiers in the geologic past. Calcifying marine organisms also would have had more time to adapt to tectonically induced acidification that occurred over geologic timescales, rather than over the human timescales that the current anthropogenic changes are occurring.

We don't know at what carbon dioxide level marine calcifiers — and their shells — will begin to disappear. Our experimental evidence suggests that this depends on the type of organism and its

mechanism of calcification. Four of the organisms we looked at exhibited a particularly rapid decline in calcification once a critical carbon dioxide level was reached: temperate corals, pencil urchins, quahogs and bay scallops. A primary objective of future studies should be to precisely constrain these apparent tipping points. The policies and legislation that govern carbon emissions should be guided by these empirically determined thresholds.

Our inability to accurately predict how marine calcifiers will respond to carbon dioxide-induced ocean acidification ultimately stems from our relatively poor understanding of the very mechanisms by which these organisms build their shells and skeletons. Much work remains to be done on this front — and others — in order to anticipate the effects of rising atmospheric carbon dioxide and guide the policy and legislation that will hopefully curtail the looming calcification crisis.

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